

Effect of Nitrogen Addition on Soil Respiration in a Larch Plantation

Jiang Zhu, Fengfeng Kang, Jing Chen, Xiaoqin Cheng, Hairong Han*

Forestry College, Beijing Forestry University, Beijing 100083, China

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Abstract

Knowledge about the impact of nitrogen (N) addition on soil respiration (R_s) is critical for understanding soil carbon (C) balance and its responses to climate change. We conducted a long-term field experiment to evaluate the response of R_s to N addition in a larch (*Larix principis-rupprechtii*) plantation during the growing season in northern China. We applied four N (in the form of NH_4NO_3) levels, i.e., control (N0: 0 g N m^{-2} year $^{-1}$), low-N (N1: 5 g N m^{-2} year $^{-1}$), medium-N (N2: 10 g N m^{-2} year $^{-1}$), and high-N (N3: 15 g N m^{-2} year $^{-1}$) monthly starting in June 2015. The results showed that N addition – especially N2 and N3 levels – significantly stimulated R_s . N1 and N2 levels resulted in the increase of R_s in a short time, whereas the duration of N3 level can last across the whole month or at least 10 days. The results highlight the need for improving the R_s sampling interval after N addition to ensure more accurate evaluation of C emission. Soil temperature and soil moisture together explained more variations of R_s . N addition exhibited a slight increase in the sensitivity of R_s to temperature (Q_{10}), but no significant differences were found for the Q_{10} among N levels.

Keywords: nitrogen supply, soil respiration, sample interval, Q_{10} , *Larix principis-rupprechtii*

Introduction

Atmospheric nitrogen (N) deposition has already increased by three- to five-fold over the past century due to fossil fuel combustion, fertilizer production, and cultivation [1]. N deposition is predicted to increase in many parts of the world in the next decades because of the rapid development and expansion of industrial and agricultural activities and thus has been considered a major scientific challenge [2]. The evidence is mounting that increased N deposition has inevitably altered regional and global environments and

has affected the carbon (C) cycle of terrestrial ecosystems [3-4].

Soil respiration (R_s) is one of the largest sources of carbon dioxide (CO_2) returning from soil to the atmosphere [3, 5]. Global CO_2 flux through soil respiration is approximately 80-98 Pg C yr $^{-1}$ [6], and annual CO_2 efflux from soil respiration is about 10 times the annual CO_2 emissions from burning fossil fuels [7]. Thus, knowledge of the response of R_s to N deposition is important for improving our ability in modeling the terrestrial C cycle [8]. However, the effect of N addition on R_s remains inconsistent for results from field observations [9], models [10], and meta-analyses [11], including stimulation [12-13], inhibition [14], and no change [15]. This inconsistency is largely related to differences in biome types, N addition rates, soil conditions, and even experimental

*e-mail: hanhr@bjfu.edu.cn

methods [16-17], The underlying mechanisms for these disagreements, however, are still unclear and need to be further investigated. The diverse responses of R_s to N addition indicates that more observations are required to understand regulatory mechanisms of N addition to the key C process in terrestrial ecosystems.

Numerous studies have investigated the effect of N addition on R_s . Generally, a large number of field and laboratory studies have examined the effect of N addition on R_s by periodic measurements of the respiratory processes. Measurements of R_s after N addition varied across studies, and the majority of sample interval from one week to one month after N addition [18-20]. However, an examination of how sampling interval affects R_s and investigation into the duration of this effect after N addition received little attention. N addition would have a pulse effect on R_s in a short time after N application. Currently, the duration of this effect after N addition is still not well known. Resolution of this question would be helpful for providing an explanation as to why R_s varies so much in response to N addition.

We used a long-term N addition experiment to assess R_s continuously and frequently after N application during the growing season in a larch (*Larix principis-rupprechtii*) plantation. Our main objectives were: a) to explore the variations in the response of R_s to N addition, b) to determine the duration of N addition effect on R_s after N application, and c) to investigate the driving factors that regulate the response of R_s to N addition.

Materials and Methods

Study Site

Our study site was located in Haodifang forestry (36°47'04"N and 112°00'57"E; 2276m a.s.l.) in Shanxi Province in northern China. This region was classified as a warm-temperate semi-arid continental, mountainous climate affected by monsoons, with a mean annual temperature of 6.2°C. The highest monthly average temperature of 22.4°C occurred in July, whereas the lowest monthly average temperature of -4.6°C occurred in January. Mean annual precipitation was 770 mm with

an average relative humidity of 65%. Over a year, the distribution of precipitation was relatively uneven. The wet season was from June to September, which accounted for more than 70% of annual precipitation (Fig. 1). Soils are developed from well-weathered rocks and are classified as a haplic luvisols (according to the FAO soil taxonomy), and it has an average depth of 60 cm. The soil organic matter content ranges from 1 to 4% w/w.

The study site was a pure plantation of *Larix principis-rupprechtii*, established with 5-year-old seedlings in 1980 at a stocking density of 1,000 stems per ha. The density of the forest now is 618 stems/ha with a total basal area of 5.68 m²/ha. The mean height of overstory was 22.5 m with a mean breast-height diameter of 18.6 cm. The understory trees were primarily *Spiraea salicifolia*, *Lespedeza bicolor*, *Corylus mandshurica*, and *Rosa xanthina*. The mean height of the understory was 1.8 m with a mean breast-height diameter of 2.0 cm. *Carex tristachya* and various other species of the family *Cyperaceae* were common herbs.

Experimental Design

At the study site, we established a random block design with three blocks as replicates. In each block, we randomly established 2 × 2 m plots (4 N addition treatments × 3 replicates = 12 plots) in April 2015. Each plot in a block was surrounded by a 2-m wide buffer strip. Each block was then assigned randomly to one of the four N treatment levels: control (N0), low-N (N1), medium-N (N2), and high-N (N3). In each plot, PVC collars (20 cm in diameter and 10 cm in height) were placed and inserted 5-6 cm into the soil for measurement of soil respiration (R_s). The location of the collars did not change during R_s measurements.

The first N addition was applied on 31 May 2015. Nitrogen was added in the form of ammonium nitrate (NH₄NO₃, AR) in water at the rate of 5 g N m⁻² year⁻¹, 10 g N m⁻² year⁻¹, and 15 g N m⁻² year⁻¹ in N1, N2, and N3 treatments, respectively. In each plot, the fertilizer was weighed, mixed with five litres of water, and sprayed to the plots using a portable sprayer at the end of each month from June to September. The N0 plots received five litres of water without fertilizer.

Soil Respiration and Microclimate

Soil respiration measurements were initiated 24 hours after the installing the collars. The R_s was measured after each N addition application from June to September in 2015. We measured R_s using an automated soil CO₂ flux system (LI-8100, LI-COR, Nebraska, USA) from 08:00 to 18:00 h, the R_s was monitored three consecutive days after each N addition application and measured 3-5 times in the remaining days of each month. The rates for each N treatment were averaged for the following data analysis.

The cumulative soil CO₂ efflux from the plots was calculated as the mean of the cumulative CO₂ effluxes of

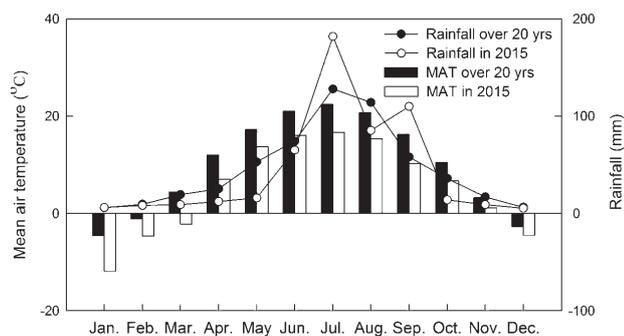


Fig. 1. Mean air temperature and rainfall during the last 20 years and in 2015.

the three replicates for each N treatment. Accumulative daily mean soil respiration efflux was calculated as follows:

$$DRs = (Rs \times 24 \times 3600 \times 12) \times 10^{-6} \quad (1)$$

...where DRs is the accumulative daily means soil respiration ($\text{g C}\cdot\text{m}^{-2}$), Rs is the daily mean soil respiration rate ($\mu\text{mol}\cdot\text{m}^{-2}\cdot\text{s}^{-1}$) in each plot, 24 is the number of hours per day, 3,600 the number of seconds per hour, and 12 represents the molar mass of carbon.

Soil temperature (ST) at 10 cm depth were measured adjacent to each Rs collar with a portable temperature probe provided with the LI-8100 system simultaneously with each Rs measurement. Soil moisture (SW) at 10 cm was measured using a moisture probe (Delta-T Devices, Cambridge, UK) at three points close to each collar for each Rs measurement.

Soil Properties and Microbial Biomass

Before the first N addition application, five soil samples were collected from the top 20 cm of the mineral soil next to each collar using 100-ml sampling cylinders in May 2015. We mixed five soil samples for soil bulk density, pH, soil organic carbon (SOC), total N (TN), total P (TP), and total K (TK) measurement. The combined samples were passed through a 2 mm sieve to remove roots and coarse fragments for further analysis. Soil pH was measured with a glass electrode in a 1:2 soil-to-water ratio. TN was determined by Kjeldahl digestion with a salicylic acid modification [21]; soil bulk density, TP, and TK were measured using standard methods [22]; and SOC using the method described by Kalembasa and Jenkinson [23]. Table 1 shows detailed soil properties for each N addition plot.

Three soil cores (10 cm in depth and 2.5 cm in diameter) were randomly collected from the top 10 cm soil from each plot in October 2015 at the end of the growing season and mixed to form a composite sample. After removing roots and plant residues, these composite samples were immediately sieved through a 2 mm mesh sieve in the field and refrigerated. Soil microbial biomass carbon (MBC) and microbial biomass nitrogen (MBN) were measured using the chloroform direct-fumigation extraction method [24].

Modeling Rs with ST and SW

The univariate and bivariate models were used to describe the relationships between Rs and ST and SW [25]. The first model shows Rs as an exponential function with only ST (ST model) as the variable (Eq. (2)), while temperature sensitivity (Q_{10}) was calculated using Eq. (3):

$$Rs = \beta_0 e^{\beta_1 ST} \quad (2)$$

$$Q_{10} = e^{10\beta_1} \quad (3)$$

...where Rs ($\mu\text{mol m}^{-2} \text{s}^{-1}$) is soil respiration, ST ($^{\circ}\text{C}$) is soil temperature at 10 cm depth, and β_0 and β_1 are model parameters.

The quadratic model was used to describe the relationship between Rs and SW (SW model) (Eq. (4)):

$$Rs = \beta_2 + \beta_3 SM + \beta_4 SM^2 \quad (4)$$

...where SM (%) is soil moisture at 10 cm depth and β_2 , β_3 , and β_4 are model parameters.

A bivariate model was used to explore the relationship between Rs soil temperature (ST) and soil moisture (SW) ($ST\&SW$ model) (Eq. (5)):

$$Rs = \beta_5 e^{\beta_6 ST} SM^{\beta_7} \quad (5)$$

...where β_5 , β_6 , and β_7 are model parameters.

Statistical Analyses

A two-way analysis of variance was used to investigate the effects of month (M) and N treatment level (N) on soil temperature, soil moisture, and Rs . For each month, repeated-measurement analyses of variance were applied to assess the effects of sample date (D) and N treatment level (N) on soil temperature, soil moisture, and Rs . One-way analysis of variance was applied to test the effect of N addition on soil properties and soil microbial biomass. The significance level was set at 0.05 ($p < 0.05$), and all statistical tests were performed using SPSS (ver. 16.0) and R3.3.2 (www.R-project.org).

Table 1. Soil properties to a depth of 20 cm in each plot of N addition treatment.

Treatment	Bulk density (g cm^{-3})	pH	Soil organic carbon (g kg^{-1})	Total N (g kg^{-1})	Total P (g kg^{-1})	Total K (g kg^{-1})
N0	0.96±0.04 a	7.6±0.1 a	20.0±0.04 a	2.32±0.03 a	0.54±0.02 a	14.9±0.0 a
N1	1.12±0.09 b	7.7±0.1 a	19.8±0.1 a	2.35±0.02 a	0.53±0.01 a	15.6±0.0 ab
N2	1.15±0.05 b	7.7±0.1 a	19.9±0.1 a	2.38±0.03 a	0.58±0.02 a	15.48±0.0 ab
N3	1.10±0.07 b	7.7±0.1 a	20.0±0.1 a	2.47±0.04 a	0.66±0.03 b	16.1±0.1 b

Data are means ± S.E. Different letters at the same column represent significant differences at $p < 0.05$.

Table 2. Results (*F*-values) of two-way ANOVA for soil temperature and soil moisture at 10 cm depth, and soil respiration rate from June to September 2015.

Source of variation	Soil temperature at 10 cm depth	Soil moisture at 10 cm depth	Soil respiration rate
Month (<i>M</i>)	4416***	12.18***	10.13***
N treatment level (<i>N</i>)	0.58	0.86	2.22***
<i>N</i> × <i>M</i>	0.34	0.06	0.05

****p*<0.001

Results

Soil Temperature and Soil Moisture

The ANOVA showed that soil temperature and soil moisture had significant monthly variation (Table 2). The highest temperature was in July and the lowest in September, and the highest moisture was observed in August and the lowest in June. N addition appeared to have no significant effects on soil temperature and soil moisture (*p*>0.05), although both had significant inter-monthly variations (Table 3).

Seasonal Variation of *R*_s

*R*_s showed a distinct seasonality, with the highest value observed in August and the lowest observed in September (Table 1, Fig. 2). The average rate of *R*_s was 3.00 μmol m⁻² s⁻¹ in June. The respective values were 3.86 μmol m⁻² s⁻¹ in July, 4.38 μmol m⁻² s⁻¹ in August, and 2.09 μmol m⁻² s⁻¹ in September. During the entire observation period, the average *R*_s rate was 3.33 μmol m⁻² s⁻¹ irrespective of the N addition effect.

N addition significantly increased *R*_s (*p*<0.05), and the rate increased with the level of N addition. The average rate of *R*_s at the N0, N1, N2, and N3 level was 2.61, 2.79, 3.30, and 4.47 μmol m⁻² s⁻¹, respectively.

*R*_s increased with N levels after N addition in each month (Fig. 2). N addition showed a strong stimulatory effect on *R*_s during the first 1-3 days. After this, the stimulatory effect at the N1 and N2 levels weakened. However, at the N3 level, N addition presented the stimulatory effect across the whole month in June, July,

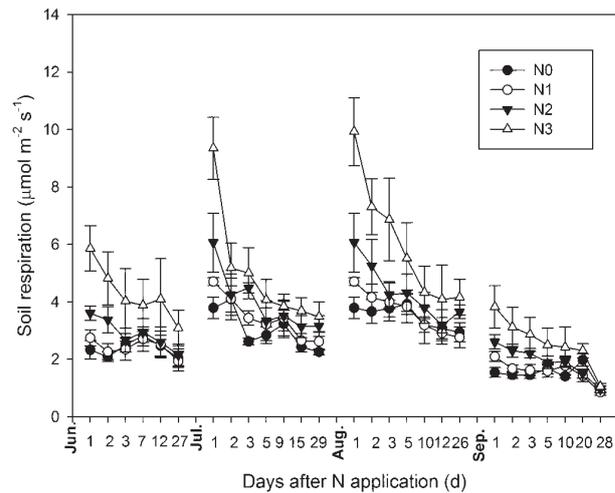


Fig. 2. Seasonal dynamics of *R*_s among N treatment levels. Error bars are ± S.E.

and August, and this effect lasted at least 10 days into September.

Cumulative *R*_s Among N Treatment Levels

The cumulative *R*_s increased with N levels after N addition in each month (Fig. 3). During the entire study period, the N2 and N3 levels significantly stimulated the cumulative *R*_s (*p*<0.05). By the end of our study, the cumulative *R*_s increased by 2.10%, 16.19%, and 51.75% at the N1, N2, and N3 levels, respectively, relative to the cumulative *R*_s at the N0 level.

Table 3. Results (*F*-values) of repeated measures for soil temperature and soil moisture at 10 cm depth and soil respiration rate in June, July, August, and September 2015. Date is the sampling date within each month.

Source of variation	Soil temperature at 10 cm depth (<i>ST</i>)				Soil moisture at 10 cm depth (<i>SW</i>)				Soil respiration rate (<i>R</i> _s)			
	June	July	August	September	June	July	August	September	June	July	August	September
Date (<i>D</i>)	212.3***	1234***	480.9***	1707***	201.20***	20.02***	8.67***	9.12***	30.09***	81.90***	16.34***	101.39***
N treatment level(<i>N</i>)	0.439	0.48	0.28	0.658	2.02	2.77	1.18	1.30	2.06	1.45*	1.36	6.21***
<i>D</i> × <i>N</i>	0.35	1.18	1.04	0.60	0.83	0.95	0.71	0.88	2.51*	8.49***	4.90***	4.54**

****p*<0.001, ***p*<0.01, **p*<0.05

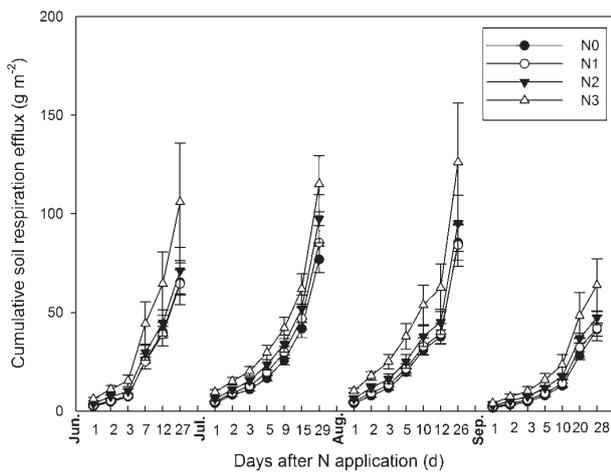


Fig. 3. Cumulative R_s efflux at different N treatment levels.

Relationships between R_s and ST and SW

Soil respiration significantly correlated with soil temperature in the four N treatments during the entire observation period ($p < 0.001$, Fig. 4). Soil temperature

explained 39.0-69.4% of the seasonal variation in R_s over the sampling period. The strongest relationship between R_s and soil temperature was found in N1 (Table 4). Estimates of the model parameter (β_1) showed that the relationship between R_s and soil temperature was not affected by N addition. Although temperature sensitivity (Q_{10}) increased in the order of $N0 < N3 < N2 < N1$, no significant differences were found for the Q_{10} among N treatments ($p > 0.05$).

There was a significant correlation between R_s and soil moisture in the four N treatments over the sampling period ($p < 0.05$, Table 3). Soil moisture explained 22.3-36.4% of the seasonal variation in the R_s .

$ST&SW$ model predicted R_s rather well, and R_s was highly dependent on the combined effects of soil temperature and soil moisture (Table 4). The $ST&SW$ model for temperature and moisture explained 72.4%, 76.2%, 70.3%, and 51.8% of the variations in R_s for N0, N1, N2, and N3 treatments, respectively. R_s in the N0 plots had the same sensitivity to soil temperature and soil moisture, whereas R_s in the N1, N2, and N3 plots was more responsive to soil temperature (Fig. 5).

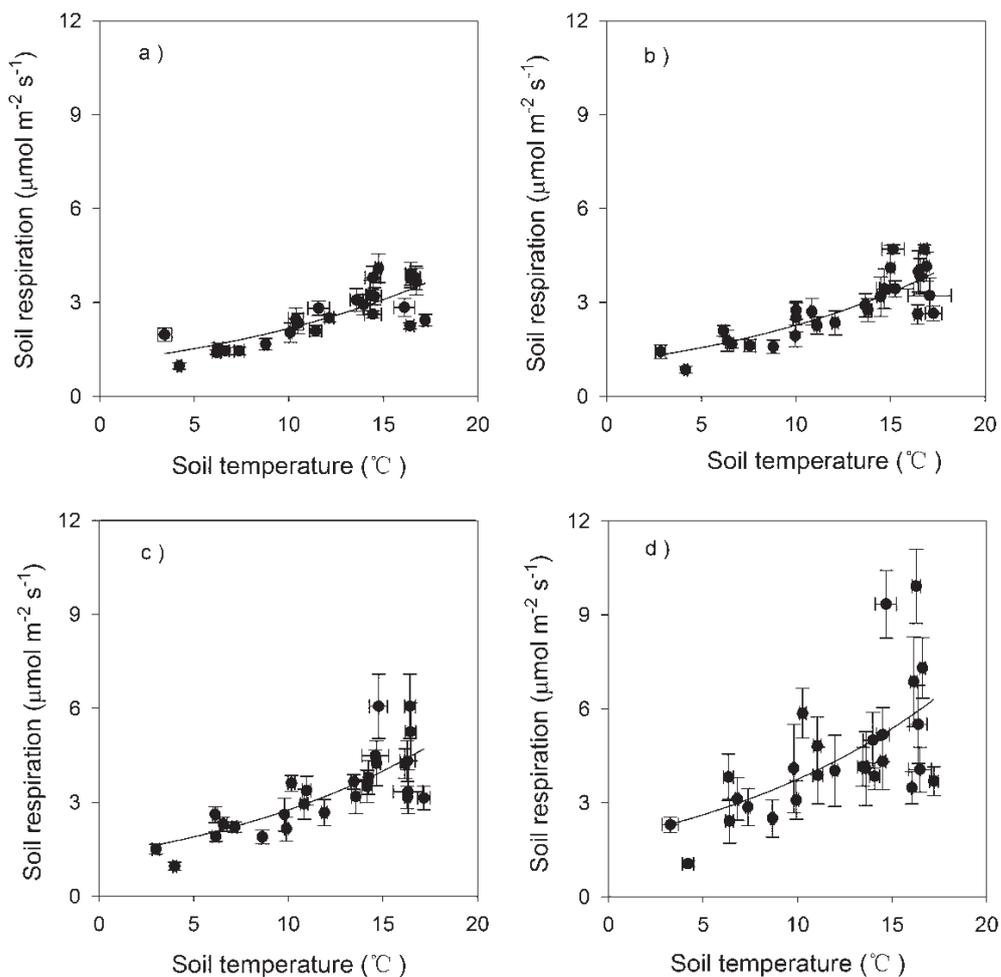


Fig. 4. Relationship between R_s and soil temperature at 10 cm depth in control (a), low-N (b), medium-N (c), and high-N (d) during observational period (June-September 2015). Points are daily means \pm S.E. The fit of the ST model (Eq. (2)) is given in Table 4.

Table 4. Parameters of different models of R_s as a function of ST and SW . Data are mean \pm S.E

N treatment	Parameters				
<i>ST</i> model	β_0	β_1	p	R^2	Q_{10}
N0	1.074 \pm 0.178	0.071 \pm 0.012	<0.001	0.662	2.03
N1	1.070 \pm 0.183	0.076 \pm 0.012	<0.001	0.694	2.14
N2	1.304 \pm 0.264	0.075 \pm 0.014	<0.001	0.602	2.12
N3	1.826 \pm 0.537	0.072 \pm 0.021	<0.001	0.390	2.05
<i>SW</i> model	β_2	β_3	β_4	p	R^2
N0	8.623 \pm 3.126	-0.489 \pm 0.205	0.009 \pm 0.003	<0.01	0.364
N1	19.481 \pm 7.479	-1.155 \pm 0.465	0.019 \pm 0.007	<0.01	0.323
N2	21.125 \pm 10.676	-1.196 \pm 0.647	0.019 \pm 0.010	<0.05	0.223
N3	26.436 \pm 8.509	-1.628 \pm 0.569	0.028 \pm 0.009	<0.001	0.318
<i>ST</i> & <i>SW</i> model	β_5	β_6	β_7	p	R^2
N0	0.857 \pm 0.243	0.074 \pm 0.010	0.004 \pm 0.007	<0.001	0.724
N1	0.924 \pm 0.328	0.081 \pm 0.010	0.002 \pm 0.009	<0.001	0.762
N2	1.112 \pm 0.363	0.082 \pm 0.016	0.001 \pm 0.010	<0.001	0.703
N3	1.687 \pm 0.407	0.080 \pm 0.017	-0.002 \pm 0.012	<0.001	0.518

Soil Microbial Biomass

The variations in MBC, MBN, and the MBC/MBN ratios at different N levels are shown in Fig. 6. The MBC and MBN increased with N levels. The MBC values at the

N3 level were significantly higher than the values at the N0 level ($p < 0.05$). The MBC at the N1, N2, and N3 levels increased by 13.4%, 37.0%, and 64.3%, respectively, relative to the N0 level. No significant differences were found in the MBN among N levels ($p > 0.05$). The

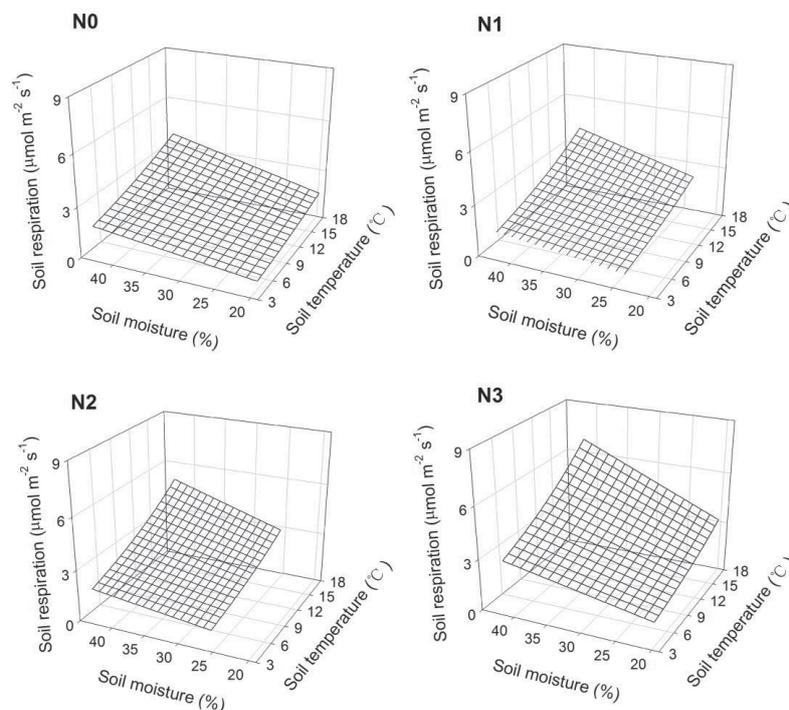


Fig. 5. 3D images of soil respiration models for N treatments with soil temperature and soil moisture. The mesh is fitted using the *ST* & *SW* model in Table 4.

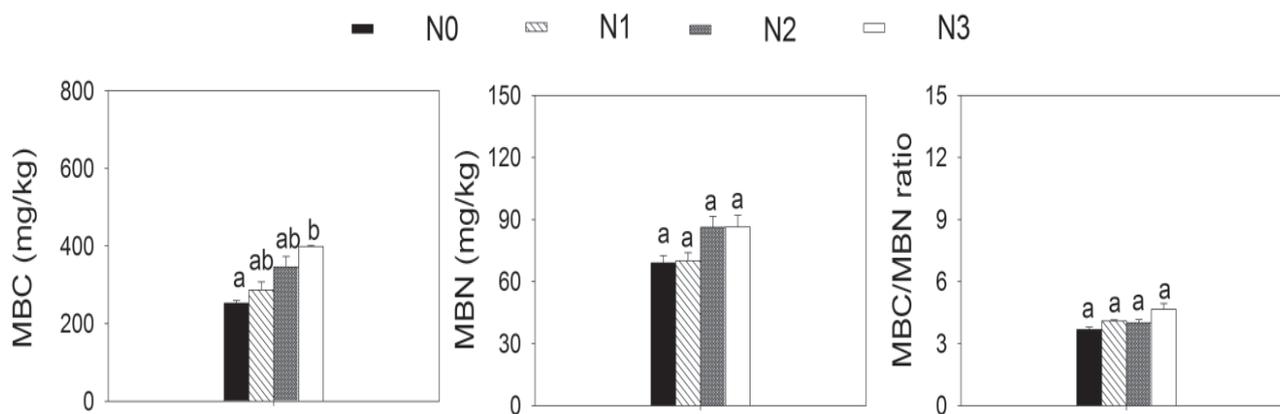


Fig. 6. Dynamics of soil microbial C (MBC), microbial N (MBN), and the MBC/MBN ratio at different N treatment levels.

MBC/MBN ratio at the N0, N1, N2, and N3 levels was 3.68, 4.09, 4.01, and 4.86, respectively. No significant differences were found for the MBC/MBN ratio among N levels ($p > 0.05$).

Discussion

Effect of Nitrogen Addition on Phytoclimate and Soil Respiration

A previous study has shown that increased plant growth and productivity under fertilization can result in shadowing, and thus a decrease in soil temperature in a temperate steppe [26]. Furthermore, a previous study also found that N-induced plant growth can result in higher evapotranspiration and more depletion of soil water content in grassland [27]. However, N addition appeared to have no significant effects on soil temperature and moisture in our study, which is consistent with results reported in a mature tropical forest [14]. This could be ascribed to the fact that forest canopy closure did not change throughout the whole observation period and the forest canopy could mediate the response of soil environment to N addition [28]. A number of studies have reported that N addition has both increased [29] and decreased R_s [30-31]. These inconsistent results could be largely due to the differential responses of two soil components: autotrophic respiration (R_a) and heterotrophic respiration (R_h). The individual response of each component to N addition determined the direction of the total R_s response [32]. N addition in the present study resulted in an increase in the R_s throughout the observation period, similar to the results of previous studies in different ecosystems [29, 33-34]. The effect of N addition on the R_s was suggested to correlate with soil background N level of studied forest. Due to N deficiency in such a young forest, external N addition could either enhance N immobilization of the microbes or improve N absorption of the roots [29]. The stimulation of R_s by N addition has been explained by N-induced increases in C substrate for root and fine root biomass, and R_s generally has a positive

correlation with the fine root biomass due to respiration from coarse root being relatively low [35]. Therefore, we speculated that N addition stimulated R_a via increasing the fine root biomass or respiration of autotrophic bacteria by root exudates [36].

Soil R_h is another part of R_s and is derived from the decomposition of aboveground litter and soil organic matter [11]. Aboveground litter respiration is dependent on the accumulated litter layers and the rate of litter decomposition. Higher net primary productivity under fertilization leads to more labile C substrate inputs into the soil through litterfall [37]. Fresh C input provides more energy to sustain the activity of microorganisms, which is typically C-limited, and thus accelerates the decomposition of recalcitrant soil organic matter and subsequently increases R_s [38]. We found that the aboveground litterfall in the study area increased significantly [39]. Similarly, Li et al. [40] studied C cycling in a tropical rain forest in Puerto Rico and found that seven years of continual fertilization increased aboveground litterfall production and stimulated litter decomposition. Magill et al. [41] found that 15 years of chronic N additions increased aboveground litterfall in a red pine (*Pinus resinosa*) plantation. A number of other studies also reported that experimental N addition increased aboveground litter decomposition by increasing litter substrate quality, which also enhanced soil microbial activities [42-43]. Therefore, the increased CO_2 efflux from the decomposition of aboveground litter and soil organic carbon under N addition was likely caused by N-induced increasing aboveground litter input, stimulating the rate of litter decomposition and enhancing soil microbial activity in the early stage of decomposition [44]. Some previous studies also have reported that N addition would inhibit R_s by decreasing fine root biomass, litter decomposition rate, and microbial activity when the ecosystem is subjected to long-term N addition and N supply exceeds plant needs. Further investigations, therefore, are needed to explore the long-term response of R_s to N addition in the studied forest. This could be helpful for enhancing our understanding of climate-C feedback of this ecosystem.

We found that the magnitude of N addition effect was seasonal, with the largest effect occurring during

June–August, which is consistent with results reported in a Chinese pine forest [45]. The duration of N addition effect was diverse among N levels. N addition presented a strong stimulatory effect on the R_s during the first 1–3 days. After this, the stimulatory effect at the N1 and N2 levels weakened. However, at the N3 level, N addition presented the strong stimulatory effect across the whole month in June, July, and August, and this effect can last at least 10 days into September. The mechanism for explaining these results might be related to the buffering aboveground litter layer and slow release of nutrients from decomposing litter added to the N addition effect.

Soil Respiration in Relation to Soil Temperature and Moisture

Soil temperature and soil moisture have been reported to directly alter the activities of plant roots and soil microbes and indirectly affect pedofauna activity, and thus alter R_a and R_h [39]. In this study, we also found that R_s exhibited significant positive relationships with soil temperature, which is in agreement with results reported in temperate forests [29] and in grassland [32]. Our model showed that soil temperature was more closely correlated with R_s than was soil moisture, both in the N addition and the control plots. This finding is also compatible with other studies [46–47]. Indeed, soil temperature and soil moisture together have been shown to play a critical role in explaining more variations of R_s than the ST or SW models alone. The Q_{10} value was 2.03 in the control plot, which is close to the median value reported in the review by Raich and Schlesinger [48], and lower than that of 3.9 in a temperate hardwood forest [49]. N addition led to a slight increase in Q_{10} in the present study, although there were no significant differences in the Q_{10} among treatments. It has been suggested that changes in the Q_{10} by N addition result from concurrent variations in soil temperature and soil water content, which may affect root and microbial activities [50]. Thus, the non-significant differences in the Q_{10} are largely attribute to the fact that there were no treatment effects on soil temperature and moisture by N addition.

Conclusions

Our data support that N addition stimulates soil respiration in an N-limited plantation. In this study, increases in R_s depend on N addition levels. N2 and N3 levels significantly increased R_s in the view of seasonal cumulative flux. The duration of N addition effect was diverse among N levels. N addition showed a strong stimulatory effect on R_s during the first 1–3 days. After this, the stimulatory effect at the N1 and N2 levels weakened. However, N3 levels presented a strong stimulatory effect across the whole month or at least 10 days. This finding highlights the need to increase the frequency of R_s measurements to capture the real R_s responses to N addition. Increasing sampling frequency

of R_s will help to understand why R_s varied so much to N addition and improve our ability in modeling C cycling in forest ecosystems. R_s was regulated by soil temperature and soil moisture and integrating soil temperature and soil moisture explained more variations of R_s . We found that N addition did not change the sensitivity of R_s to temperature throughout the sampling period.

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