

Original Research

Estimating the Effect of Precipitation and Vegetation on Water Yield in Northern China

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Abstract

This study attempts to analyze how precipitation and vegetation affect water yield in Haihe watershed. Based on the annual runoff and meteorological data, the water balance equation and the calibrated Zhang model were used to simulate the effects of precipitation and vegetation on water yield. The simulated results showed that the effect of precipitation on water yield was linear in mountainous area of Haihe watershed, and a 1 mm change (increase or decrease) in precipitation will cause a 0.51 mm change in water yield; under the mean annual precipitation of 534 mm, the effects of decreasing water yield by per unit cover ratio for forest, grassland, and farmland were 1.26, 0.47, and 0.69 mm, respectively. The decreased water yields for forest, grassland, and farmland were associated with aridity indexes by negative power functions, and indicate that the effect of vegetation on water yield was smaller in drier areas. For three vegetation types the effects of vegetation on water yield were similar for crops and grass, and were much larger for forest.

Keywords: precipitation, vegetation, water yield, northern China

Introduction

The effect of vegetation on water yield is a popular and important issue in ecohydrological studies. In response to the widely observed degradation of formerly forested land and the rising demands for paper pulp, industrial wood, and wood fuel, the need for large-scale reforestation programs has been expressed repeatedly. To address the increasing levels of soil erosion and desertification, China has implemented many large-scale reforestation programs in recent decades [1]. While creating many economic and ecological benefits, reforestation could also

potentially reduce water capacity [2-3]. Because of this, correct understanding of the forest-water relationship is particularly important.

There have been many influential studies made in understanding forest and water relationships during the past century around the globe. Our understanding of the forest impact on annual water yield is well advanced and there are robust methods available for predicting the impact of forest change on the mean annual water balance [4]. Methods have also been established that allow the prediction of water yield changes in response to forest change at the annual time scale [5-6].

Although researchers concurred with that forests decrease water yield, the decreasing extent is different. Farley et al. (2005) made a global analysis using 26

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catchment datasets and 504 observations, found that the annual water yield was reduced by 44% and 31% when grasslands and shrublands, respectively, were afforested [7]. Sun et al. (2006) examined the sensitivity of water yield responding to forestation across China, then recorded that the reduction in water yield due to forestation varied from about 50 mm/yr (50%) in the semi-arid Loess Plateau in mountainous area of Haihe watershed to about 300 mm/yr (30%) in the tropical southern region [8]. Wang et al. (2011), using published literature data from the past 50 years, showed that the regional average of annual runoff for forest is only 16 mm – 58% lower than the 39 mm for non-forestland [9]. These findings indicate that the effects of vegetation can vary with climate and/or geography.

Evapotranspiration is the main process responsible for changes in mean annual water yield in response to alterations in vegetation [10-12]. Zhang et al. (2001) analyzed results from 250 catchments worldwide [13]. They developed a simple two-parameter model to estimate mean annual evapotranspiration at catchment scale. This model allowed for simulation of the effects of precipitation and different vegetation.

The Haihe watershed is the largest water system in northern China and covers the fastest growing economic region of China [14]. This region is also densely populated, highly developed, and the most water-scarce region in China. It has been transferring real water from outside through various water diversion projects [15]. In recent years, the ecological environment of the watershed has been severely damaged as a shortage of water resources and a series of ecological environmental problems have made the region the most prominent contradiction between water and humanities in China [16]. A study of changes in water yield in this region thus has an important practical significance.

The objectives of our study were to:

1. Analyze the quantitative effect of precipitation on water yield.
2. Estimate the effects of forestland, grassland, and farmland on decreasing water yield.
3. Investigate the effects of vegetation on water yield with consideration of aridity indices.

Material and Methods

Study Area

The study was carried out in the mountainous areas in Haihe watershed, which is bound by latitude 35°4' and 44°11' and longitude 111°41' and 120°17' (Fig. 1). The areal coverage of watershed is 31.82×10⁴ km² and selected mountainous area 26.48×10⁴ km². The topography is characterized by low mountains and hilly landscapes, with elevation varying from 1 to 2,940 m above sea level. The soil layer is 30~80 cm. The main land use types in the study area were forest, grass, and crop, accounting for 31.82%, 31.41%, and 32.06%, respectively. The principal

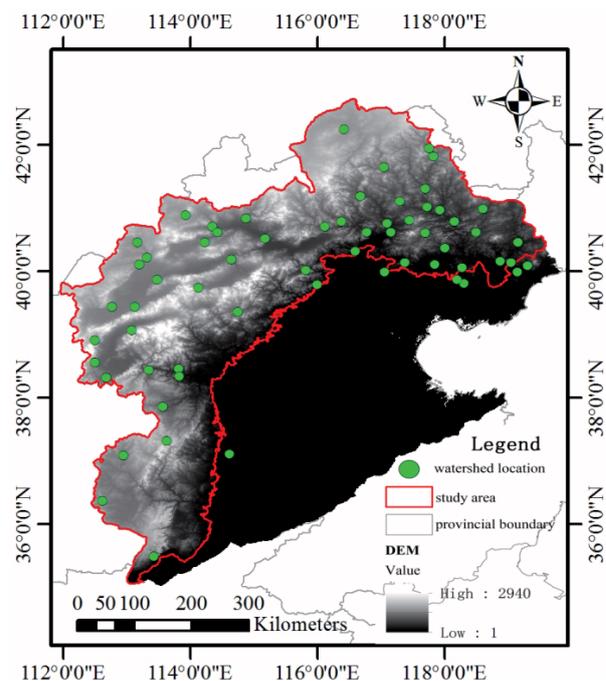


Fig. 1. Location of study area and catchment outlets.

species of forest were *Pinus tabulaeformis*, *Platycladus orientalis*, and *Quercus variabilis*; principal species of grass were *Gramineae*, *Cyperaceae*, and *Compositae*; and the principal species of crop were *Zea mays* L. and *Triticum aestivum* L.

The region has a semi-humid continental monsoon climate with a mean annual temperature of 12°C. Precipitation is highly variable, both spatially and temporally, in the range 350-800 mm. More than 85% of the rain falls between June and October.

Data

Annual precipitation and runoff were obtained from the Hydrological Yearbook of PR China (1960-91). Meteorological data were obtained from the China Meteorological Data Sharing Service System (data.cma.cn/en). The digital elevation model (DEM) of this region, with a resolution of 90×90 m, was used to delineate the catchment boundaries of all catchments. The DEM and forest cover data were both derived from Data Sharing Infrastructure of Earth System Science (www.geodata.cn).

Methods

Water balance provides a common framework for studying hydrological behavior in watershed scale, and it can be expressed as:

$$P = ET + R + D + \Delta S \quad (1)$$

...where P is mean annual precipitation (MAP) (mm), ET is mean annual evapotranspiration of a watershed (mm),

R is mean annual runoff (MAR) (mm), D is groundwater recharge, and ΔS is the change in soil moisture. D and ΔS can be considered to be 0 for a sufficiently long time series (>10 years) (Zhang et al. 2007). This formula can then be simplified as:

$$P = ET + R \tag{2}$$

By analyzing results from 250 catchments worldwide, Zhang et al. (2001) developed a simple two-parameter model to estimate mean annual evapotranspiration at the catchment scale [13]. Evapotranspiration can be described and estimated by:

$$\frac{ET}{P} = \frac{1 + \omega \frac{E_0}{P}}{1 + \omega \frac{E_0}{P} + \left(\frac{E_0}{P}\right)^{-1}} \tag{3}$$

...where E_0 (mm) is potential evapotranspiration and ω is the plant available water coefficient, which represents the relative difference in the way plants use soil water for transpiration.

Sun et al. (2005) developed this model for catchments with mixed land use [17]:

$$ET = \sum (ET_i \times f_i) \tag{4}$$

...where f_i is the percentage of land use i , including forestland, grassland, farmland, bare land, water bodies, and residential lands.

Potential evapotranspiration (E_0) is closely associated with temperature, solar radiation, and other meteorological factors, and varies widely in different watersheds. The potential evapotranspiration (E_0) is calculated as:

$$E_0 = \alpha \frac{\Delta}{\Delta + \gamma} R_n \tag{5}$$

$$\Delta = 2503 \frac{e^{\frac{17.27T}{T+237.3}}}{(T+237.3)^2} \tag{6}$$

...where α is a constant equal to 1.28, Δ is the slope of the saturation vapor pressure curve, γ is the psychrometric constant, T is the daily mean air temperature, and R_n is total net radiation.

The main land use types in mountainous area of Haihe watershed are forestland, grassland, farmland, bare land, and residential land. In simulations of the Zhang model, the three main model parameters are P , E_0 , and ω (Eq. 3). To extend the model to catchments with various land uses, Zhang et al. (2001) assumed that annual evapotranspiration of a catchment was the sum of forested and non-forested land, and simply represented non-forested land as

grassland [13]. Land use in the study area is very complex, however, so the simple representation of non-forested land as grass is not appropriate. We thus had to calibrate E_0 with Eq. 5 for the different land uses before we could use the Zhang model.

In statistics, the mean absolute error (MAE) is a quantity used to measure how close forecasts or predictions are to the eventual outcomes. The mean absolute error is given by:

$$MAE = \frac{1}{n} \sum_{i=1}^n |x_i - y_i| \tag{7}$$

...where x_i is the prediction, y_i is the true value, and n is the sample size.

Results and Discussion

Fitting Model Parameters

Potential evapotranspiration was calculated with Eq. 5 using long-term (1960-90) meteorological data for each of the 50 national weather stations within or around the study area. Among the 21 stations within the study area, 14 were surrounded by farmland, four were surrounded by grassland, and three were surrounded by forests. The potential evapotranspiration of forest, grassland, and farmland was calculated from the mean values of these stations, respectively. No weather stations were surrounded by bare land, water bodies, or residential land, so we simulated the potential evapotranspiration of the study area (Fig. 2) by Kriging interpolation

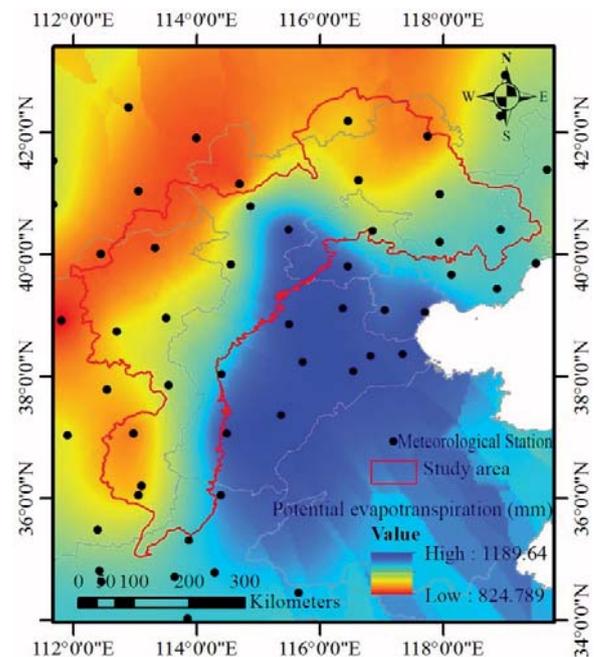


Fig. 2. Potential mean annual evapotranspiration in mountainous area of Haihe watershed.

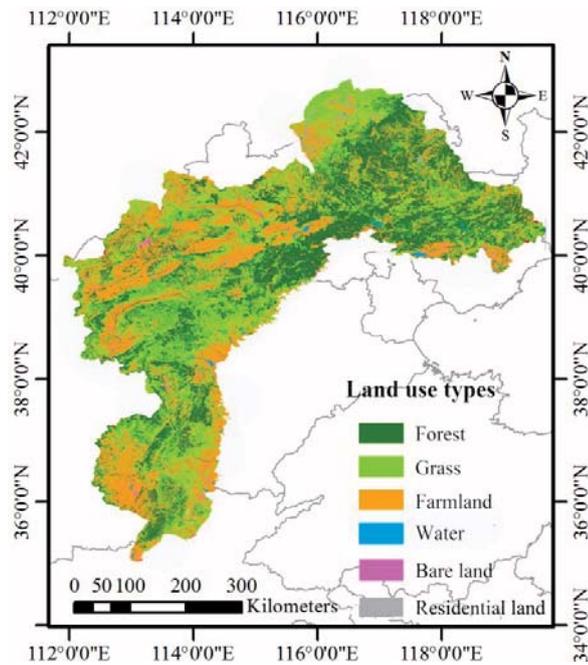


Fig. 3. Land use types of mountainous area of Haihe watershed.

using ArcGIS software, and then extracted the potential evapotranspiration for these three land use types by using the land use map (Fig. 3).

According to Zhang et al. (2001), the best fit values of ω is 2.0 for forest, 0.5 for grass and crop, and <0.5 for bare soil. To best fit the ω values to study region, we tested each group ω from 1.5 to 2.5 for forest, from 0.1 to 1.0 for grass or crop, and from 0.0 to 0.5 for bare and residential land. The best simulations were for ω values of 2.0 for forest, 0.6 for grass/crop, and 0.1 for bare/residential land. The derived values of E_0 and ω for the mountainous area of Haihe watershed are shown in Table 1.

Fig. 4 shows the comparison of measured and simulated evapotranspiration values using the calibrated parameters and the Zhang model for the 32 watersheds. Linear regression analysis indicated that the measured and simulated data were highly correlated ($R^2 = 0.89$, $P < 0.001$), the slope of the regression line is 0.98, which is close to 1; the mean absolute error (MAE) (Eq. 7) between the model estimates and the measurements was 20.84 mm, or 4.70% of the mean annual ET. Thus the calibration of

Table 1. Calibrated parameters of the Zhang model.

Land use types (<i>i</i>)	E_0 /mm	ω
Forest	1,000	2.0
Grass	960	0.6
Farmland	980	0.6
Bare land	910	0.1
Residential land	920	0.1
Water bodies	1,150	-

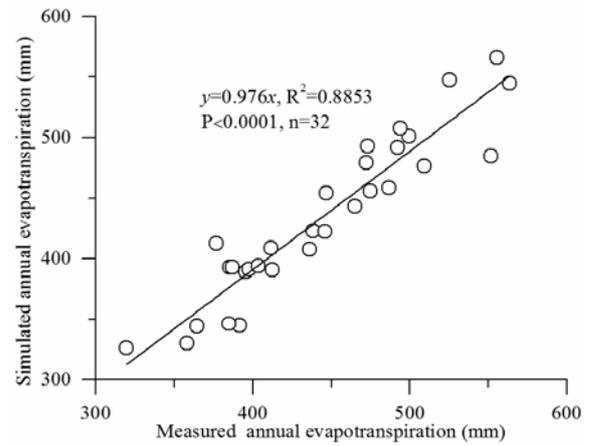


Fig. 4. Model calibration: correlations between simulated and measured annual evapotranspiration for 32 catchments.

the model parameters (E_0 and ω) was good, and these parameters could be used in subsequent simulations.

Effects of Precipitation on Water Yield

Precipitation is the main source of water yield and also the most important influencing factor. The mean annual precipitation in north China is 534 mm. In this study, based on the Zhang model (Eq. 3) and water balance equations (Eq. 2), we thus used 534 mm to simulate the corresponding water yield of each catchment. The measured and simulated values in Fig. 5 show the response of water yield to variations in actual levels of precipitation from 534 mm. In this figure, x-axis showed the difference between actual precipitation and simulated precipitation of 534 mm, and then y-axis showed the difference between the actual streamflow in actual precipitation and the simulated streamflow in simulated precipitation at 534 mm.

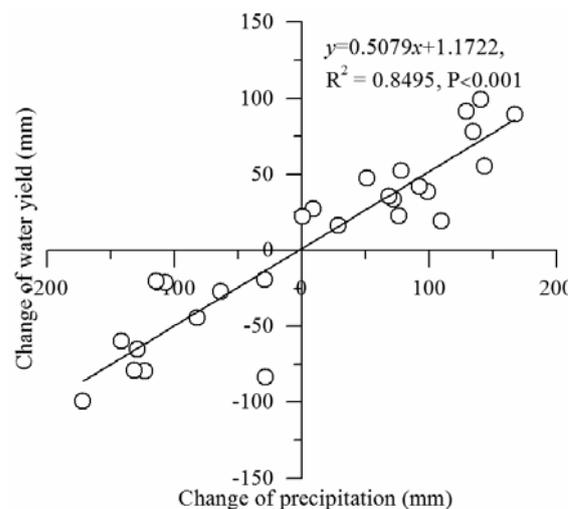


Fig. 5. Response of annual water yield caused by change of precipitation from actual to 534 mm.

Table 2. Correlations of precipitation and the water yield change of three vegetation types.

	Forest	Grass	Crop
Pearson correlation	0.600**	0.566**	0.569**
Sig.	0.001	0.002	0.002

**Correlation is significant at the 0.01 level

Fig. 5 shows the changes of water yield and precipitation with a good linear correlation (Table 2), and the equation for the regression line is $y = 0.51x + 1.17$ ($R^2 = 0.85$, $P < 0.001$). The effect of precipitation on water yield was thus linear, the water yield increases/decreases linearly with increasing/decreasing precipitation. The slope of the fitting line was 0.51, i.e. a 1 mm change (increase or decrease) in precipitation will cause about a 0.51 mm change (increase or decrease) in water yield.

Mean annual precipitation is one of the most important determinants of annual runoff and can have a strong influence on changes in runoff after a change in vegetation [18-20]. Amarasingha undertook a further review of paired catchments and concluded that changes in water yield were largest in areas with high rainfall [21]. Huxman Travis E. et al. also suggested that forests influenced evapotranspiration to a greater extent in more mesic climates and less so in arid climates [22]. In agreement with these previous analyses, our results show that the effects of three vegetation types on decreasing water yield all increased with precipitation and could be described by power functions ($Y = ax^b$). Vegetation thus has a larger effect on water yield in areas with more precipitation. The negative power relationships ($Y = ax^b$) between the effects of vegetation and the aridity index also show that vegetation has a smaller effect on water yield in drier areas.

Effects of Vegetation on Water Yield

In this study, the effects of clear-cutting the vegetation (to become bare land) on water yields have been simulated by a reduction of the ω parameter from original value of different vegetation types to 0.0, and the difference between the simulated and measured values indicated water yield's lowering effects of forest, grassland, or farmland on water yield [17]. The amount of water yield reduction was closely associated with the cover ratio of three vegetation types (Fig. 6).

The lowered water yields of these three vegetation types were well correlated linearly with their cover ratio ($R^2 > 0.70$, $P < 0.001$). The lowering effects of these vegetation types on water yield thus increased linearly with cover ratio. The slopes of the regression lines represent the lowering rate water yields (mm) of per unit percentage (1%), and we use them to reflect the sensitivity of vegetation on decreasing water yield. As seen from Fig. 6, the powers of grass and crop on decreasing water yield are approximately 0.47 and 0.69 mm/%, while the

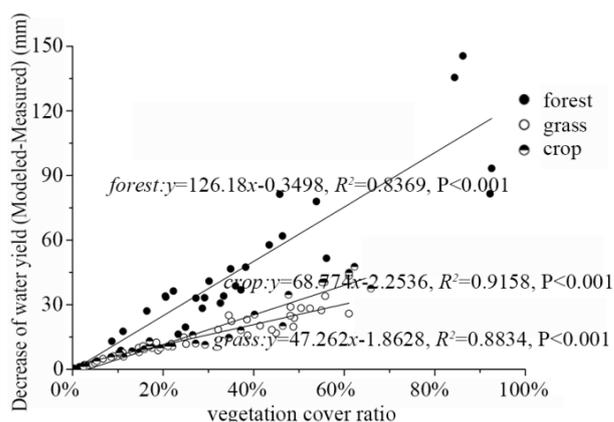


Fig. 6. Relationships between decreased water yield and vegetation cover ratio.

power of forest is 1.26 mm/%, much larger than the grass and crop. That means a 1% decrease in the cover ratio of forest, grass, and crop would increase water yield 1.26, 0.47, and 0.69 mm.

Zhao et al. illustrated that water yields changed following vegetation changes, and this process occurred not only in small catchments but also in larger watersheds [23]. In this study, of the three vegetation types (forest, grass, and crop), forest had a much larger effect on water yield than did the other two types. Forest have a higher capacity for water consumption, associated with the higher leaf area index (LAI) of the higher stature vegetation [24], and tend to have better access to water sources through accessing deep water or drawing on stored soil water [25]. Transpiration is traditionally considered the most important component of forest evapotranspiration, but interception and subsequent evaporation from the canopy can also increase substantially, particularly with conifers [26]. The evaporation of intercepted precipitation can account for 10-20% of the rainfall for broadleaf trees and 20-40% for conifers [27]. All these factors contribute to the larger effects of forests on water yield than the effects of grass and crop.

Effects of Vegetation on Water Yield Considering Aridity Indices

The aridity index, as a function of precipitation and potential evaporation (E_0/P), represents the degree of drying in a region and can be used as a good representative indicator of climatic conditions [28-29]. Fig. 7 shows the variations of the affected water yields due to vegetation change of 1% with different values of the aridity index. The results show that the lower effects of three vegetation types on water yields are all decreased with the aridity index, and could be described by negative power functions ($Y = ax^b$), indicating that vegetation has less of an effect on water yield in drier areas. The effects of crop and grass were similar, and the effect of forest was much greater.

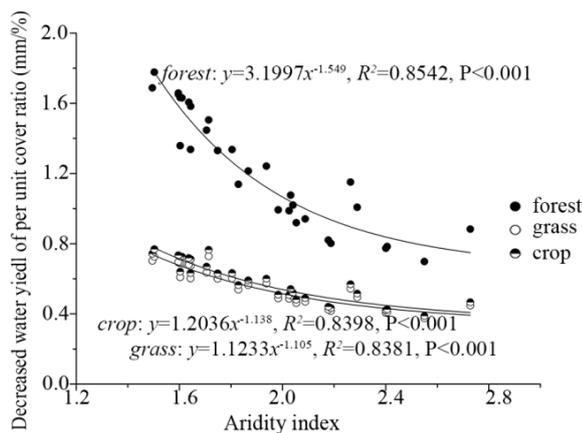


Fig. 7. Relationships between affected water yield and aridity index.

Many previous studies showed that the changes of streamflow were different in different regions and in different time scales, and especially for the effects of climate change. There was a finding of the previous studies that low-streamflow area was more sensitive to climatic variation than high-streamflow area. And the effect of vegetation, based on the Zhang model, was caused by their evapotranspiration, and had a close relationship with the climate.

Previous research indicated that the stage of forest development also affected water yield. In this study, the age of trees in the study area were mostly older than 50 years, and we also deliberately used the average values of the records of longer duration (10-31 years) in an attempt to minimize these errors, and this has already been a simple and effective approach.

In large forested watersheds, forest disturbance and climatic variability are commonly recognized as two major drivers interactively influencing hydrology in forested watersheds [30-31]. The biggest challenge is how to separate their relative contributions to hydrology [32-33]. And some hydrological modeling is often used to assess the relative effects of climate variability and forest change on hydrology [23]. But this modeling approach requires major efforts on model calibration and has uncertainties in the structure and parameters of the model, and is only applicable for the watersheds with long-term available data on vegetation, soil, topography, land use, hydrology, and climate [31]. In this study, we simply simulated the effects of vegetation on water yield with a given precipitation of the mean annual value of 534 mm, and attempt to remove the disturbance of precipitation. And this might lead to some uncertainties.

Conclusions

The impact of vegetation manipulation on water yield was one of the important components in ecohydrological studies. However, the direction and magnitude of water

yield change associated with vegetation has long been an issue of debate. Analyzing the effects of precipitation and vegetation on water yield could be helpful to understand the hydrological relationship between vegetation and water. In this study, we simulated their effects by using the water balance equation and the calibrated Zhang model and concluded that:

1. In this study area within certain precipitation, the effect of precipitation on water yield was linear, and a 1 mm change (increase or decrease) in precipitation would cause a 0.51 mm change (increase or decrease) in water yield.
2. With mean annual precipitation of 534 mm in the mountainous area of Haihe watershed, the effects of decreasing water yield by per unit cover ratio (1%) for grassland and farmland were approximately 0.47 and 0.69 mm, respectively, while the effect by per unit cover ratio (1%) for forestland was 1.26 mm – much greater than grassland and farmland.
3. Vegetation had less effect in drier (larger aridity index) area; the effects of the three vegetation types were associated with aridity index by negative power functions ($Y = ax^{-b}$); the effects of crop and grass were similar, and the effect of forest was much greater.

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