Morphological Characteristics and Soil Mechanical Properties of a Gully in the Dry-Hot Valley Region of Southwestern China

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Abstract

Gully erosion is a dominant process causing severe soil loss and soil degradation in the Dry-Hot Valley region of southwestern China. Little attention has been paid to the spatiotemporal variations of the morphological characteristics of gullies. The aims of the present study were to explore gully morphological characteristics and its soil mechanical properties. In our study, we identify the main types of gullies according to their activity and the characteristics of their impact factors in the study area. Results showed that 1) three different active gully types could be classified: active, semi-active and stable; 2) the morphological characteristics of different active gullies were significantly different; 3) the soil shear strength under different parts of gullies varies with different active degrees of gullies; and 4) the gully head-cut height, gully wall-slope and its bed-gradient ratio were the most important indicators for reflecting gully morphology, which was affected by soil mechanical properties in it. It is the premise and foundation of further study on gully erosion to carry out the study of a gully’s morphological characteristics. These findings improve the understanding of a gully’s morphological characteristics and its soil mechanical properties in the Dry-Hot Valley region.

Keywords: morphological characteristics, mechanical properties, gully, Dry-Hot Valley region

Introduction

Gully erosion was one of the serious environmental problems in the Yuanmou Dry-Hot Valley downstream of the Jinsha River [1-5], where soil erosion intensity reached up to \(1.64 \times 10^4 \text{ t} / \text{(km}^2\text{•a})\) and gully density was up to 3.0–5.0 km/km\(^2\) [3-5]. At present, there was a great deal of research focusing on one of the parts among gully head, gully wall and gully bed in the region [3-6], while the comparative studies on various parts of gullies are seldom reported. The soil shear strength was one of the measuring indicators for studying soil stability [7-10], which could characterize the stability...
of soil and explain the reason for erosion difference for different parts. For the soil mechanical properties, numerous research has focused on the influence of the cohesive forces on shear strength characteristics [11-15], and the effect of fines on the mechanical properties [16, 17]. Ouyang et al. [18] concluded that soil shear strength exerted a huge impact on soil erodibility, which thus could be used to determine soil erosion intensity. Helson et al. [19] used soil shear strength to study the mechanism of soil splashing erosion and proved that there was a close relationship between the amount of splash erosion and shear strength. But there was little research on the influence of mechanical parameters on the gully, especially in the special geomorphology of a well-developed gully region. Gully erosion could be divided into different developmental stages according to the visible phenomenon such as the amount of collapsed soil [2-6].

Long-term field surveys found that gullies at different active stages experienced obvious changes with different morphological characteristics of their headcut heights, wall-slope and bed width. Bennett et al. [20] applied fractal geometry theory to explore the gully fractal curvature characteristics. Wuddivira et al. [10] integrated GPS instruments into gully morphology monitoring. At the same time, numerous studies had shown that hypsometric integral (HI) was a kind of macroscopic topography index with definite physical and geomorphological meaning as one of the quantitative analysis parameters for the stage of geomorphologic development [21-24]. The geomorphic information entropy theory was a method derived from information entropy principle to determine the degree of topographic development [24, 25], which proved to be suitable for the quantitative study regarding the small-scale erosion topography [13]. It could be seen that it was feasible to judge the extent of gully erosion using the geomorphic information entropy theory.

Gullies are formed due to a height drop caused by a terrace or river bank, which develop by way of headward retreat on erodible hillslopes [2-6]. Furthermore, the gully slopes are often more steep upstream (compared to downstream) of gully headcut heights [1]. Studies had shown that gully morphological characteristics were an effective agent for geomorphic change in gully systems [26-29]. These studies provide important insight into processes and mechanisms of gully landform evolution and erosion. However, research on the relationship between gully activity and morphological characteristics was still rarely reported. Related research mainly focused on the qualitative study on gully evolution lacking quantitative criteria for evaluation factors. In this region, in particular, morphological characteristics of gully erosion have not been well quantified. Since gully activity dominates the development direction of the whole gully system, quantitative study on determining the different gully activities was of great significance to the controlling practice. In this study, therefore, the RTK-GPS instrument and GIS technique were applied to evaluate the activities and morphological characteristics of gully development. The soil shear strength indices for different parts of a gully were measured and the differences of mechanical properties were discussed.

**Material and Methods**

**Study Area**

Experiments were carried out at the Yuanmou Gully Erosion and Collapse Experimental Station, a field station operated by the Chengdu Institute of Mountain
Hazards and Environment (IMHE), Chinese Academy of Sciences (CAS). The station is situated within Yuanmou County (25°23′N to 26°06′ N, 101°35′ E to 102°06′E), Yunnan Province, China, and is part of the Jinsha River Basin (Fig. 1). The study area is under the influence of a dry-hot climate with a mean annual precipitation of 634.0 mm, a mean annual temperature of 21.8ºC, and an average annual potential evaporation of 3847.8 mm. Rainstorm events typically last from one to two hours and approximately six to seven major high-intensity storms occur in this area each year. Total evaporation is greater than precipitation by 5.8 times, which accounts for the dry-hot climate of the region [25]. The dominant soil type is dry red soil with bulk densities of 1.4 to 1.8 g cm\(^{-3}\), respectively, classified as Ustic Ferrisols in Chinese Taxonomy and Ferralic Cambisol in the FAO soil taxonomic system [26]. The soil erosion rates were estimated at 1.64 × 10\(^4\) t km\(^{-2}\)a\(^{-1}\) in the Yuanmou Dry-hot Valley Region [25]. Gully erosion accounts for much of the total soil erosion that takes place, having a gully density between 3 and 5 km km\(^{-2}\) [1-4]. Accordingly, this area in the mountainous region of southwestern China is considered an ecologically fragile zone. Moreover, the area is difficult to revegetate due to its extreme climate and serious soil erosion. Zonal vegetation type is tropic bushveld with a scattering of trees. The vegetation of the dry-hot valley region is dominated by herbs belonging to the Heteropogon genus and Bothriochloa pertusa species as well as sparsely dispersed Leucaena leucocephala and Dodonaea viscosa shrub species. It consequently resembles a tropical savanna ecosystem [1-4].

**Methods**

*Extracting Gully Morphological Data Based on ArcGIS Software*

Thirty-six gullies developed in the dry-hot valley were selected and measured by RTK-GPS instrument (Trimble R8 with 10 mm±1 ppm RMS dynamic measurement and 20 mm±1 ppm RMS dynamic accuracy, respectively). Using ArcGIS 10.2 software, the measured scattering coordinate data is processed and analyzed to generate a digital elevation model (DEM) with a resolution of 0.1 × 0.1m. Using DEM to extract gully morphological parameters (gully width, gully head-cut height, gully wall-slope and other parameters et al.). The specific steps were as follows (Fig. 2): 1) import RTK-GPS database into ArcGIS software as point files, 2) generate tin models using Arc Map’s 3D analyst tool, 3) convert tin to a digital elevation model (DEM) with a resolution of 0.1 m × 0.1 m using the convert tin to grid tool, and 4) extract morphological parameters.

*Application of Geomorphic Information Entropy to Judge the Activeness of Gully Development*

Hypsometric integral (\(HI\)) could be expressed as follows:

\[
HI = \frac{(H_{\text{mean}} - H_{\text{min}})}{(H_{\text{max}} - H_{\text{min}})} (1)
\]

...where \(H_{\text{mean}}\) is average elevation, \(H_{\text{min}}\) is minimum elevation, and \(H_{\text{max}}\) is maximum elevation.

The geomorphic information entropy value could determine the extent of the development of landform erosion and should be calculated as follows:

\[
H = S - \ln S = \int_0^1 f(x)dx - 1 - \ln [\int_0^1 f(x)dx] (2)
\]

...where \(H\) is the geomorphic information entropy value, \(S\) is Strahler area-elevation integration value, and \(f(x)\) is the Strahler area -elevation fitting function curve.

*Direct Shearing Tests*

The soil samples were taken from the gully head, gully wall and gully bed, respectively, in each gully of the selected 36 ones. The surface soil was collected with a 60 cm\(^3\) ring cutter (matched with the direct shearing device), the sampling depth was 2 cm, and 4 replications were set for each part. The soil samples were collected with plastic wraps and were sealed immediately. After
completing the pre-treatments, the shearing tests were carried out on the strain-controlled direct shearing device with vertical loads of 100 k Pa, 200 k Pa, 300 k Pa and 400 k Pa, respectively, and a shear rate of 0.8 mm / min. Obtained by the vertical stress and shear stress in the coordinate system, the number of points were connected with the line, that was, the destruction of the envelope, while measuring soil friction angle (φ) and cohesion (c), and could be expressed as follows:

$$\tau = \sigma \tan \phi + c \quad (5)$$

...where $\tau$ is the soil shear strength, k Pa; $\sigma$ is the normal stress acting on the shear plane, k Pa; $\phi$ is soil friction angle (°); and $c$ is cohesion, k Pa.

Data Analysis

ArcGIS10.2 was used to extract gully morphology data. All statistical analysis were carried out using Origin software (version 8.0).

Results and Discussion

Morphological Characteristics of Different Activeness Gully

Gully Classification and Spatial Distribution Trend with Different Activity

According to the research results of the topography information entropy, the geomorphic erosion could be divided into the following three different stages. When the value of $H$<0.1110, the geomorphologic development stage was the infancy stage; when 0.111 0≤$H$≤0.400 0, it was the maturity stage; when $H$>0.400 0, which was the grontic stage. Based on the above classification criterions, the gullies in the study area belonged to the maturity stage of geomorphic erosion. According to the field investigations for gully development, this paper classified the maternal age entropy (0.1110~0.4000) into three grades, namely named active period (0.1110≤$H$≤0.1885), semi-active period (0.1886≤$H$≤0.2225) and stable period (0.2226≤$H$≤0.4000). The Significant differences were found in the values of information entropy between the active, semi-active and stable gullies (Table 1).

### Table 1. Gully activity characteristics based on information entropy.

<table>
<thead>
<tr>
<th>Information entropy classification criterions</th>
<th>Activity characteristics</th>
<th>Relative height /m</th>
<th>Area/m$^2$</th>
<th>$S$</th>
<th>$H$</th>
</tr>
</thead>
<tbody>
<tr>
<td>0.1110–0.1885</td>
<td>Active gully</td>
<td>3.88±2.36a</td>
<td>58.26±54.51a</td>
<td>0.55±0.03a</td>
<td>0.15±0.02c</td>
</tr>
<tr>
<td>0.1886–0.2225</td>
<td>Semi-active gully</td>
<td>3.41±1.61a</td>
<td>42.72±42.39b</td>
<td>0.49±0.01b</td>
<td>0.21±0.01b</td>
</tr>
<tr>
<td>0.2226–0.4000</td>
<td>Stable gully</td>
<td>2.84±1.06b</td>
<td>43.67±34.58b</td>
<td>0.45±0.01c</td>
<td>0.25±0.02a</td>
</tr>
</tbody>
</table>

Note: Data in the table indicate the mean value ± the standard value, n = 12. Different lowercase letters for the same stage show significant differences among different gullies at the 0.05 level.
development, the greater the bed-gradient ratio, and the gully bed was dominated by runoff erosion. On the contrary, the more stable the gully development, the smaller the bed-gradient ratio, the weaker the water erosion, and the gully bed was dominated by sediment. Field research could verify this conclusion to some extent by the serious erosion occurring in an active gully head due to dropping water from the upstream area drop-falling erosion eroded violently in an active gully head, and in a stable gully due to long-term sedimentation, the gully bed topography tended to be flat.

Mechanical Properties of Gullies with Different Gully Activeness

Soil Shear Strength of Gullies with Different Activeness

The soil particles and aggregates were arranged in a certain form of soil structure. The frictional forces of different particles and the occlusal forces generated by the intercalation and interlocking of particles affect their ability to resist the destruction of external forces. The soil shear strength in different parts of gullies were plotted in the form of Mohr stress circle (Fig. 3). The shear strength of active gully head under 4 confining pressure showed the largest, and the active gully bed was the smallest. This might be the development of a large number of ferromanganese film in active gully head. The shear strength of semi-active gully was the smallest under the 4 confining pressures. When the pressure was less than 200 kPa, the shearing strength of the gully wall was greater than that of the gully bed. After the confining pressure increases to 200 kPa, the shearing strength of the gully bed was the highest. Field survey found that the semi-active gully head exist headward erosion and vegetation had growth on a semi-active gully bed. When the confining pressure was less than 400kPa, the shearing strength of the gully wall was smaller than that of the stable gully head. When the confining pressure exceeds 400kPa, the shearing strength of the stable gully wall was larger than that of the stable gully head. This indicates that the information entropy of stable gully reaches 0.25, soil erosion was relatively light, and almost no collapse occurs in the stable gully wall. Especially the stable gully bed had undergone many years of physical weathering, surface and subsurface biological functions, soil particle structure was good, and mechanical properties were stable.

The Internal Friction Angle of Gully Various Parts

There was a significant linear positive correlation between soil friction angle and soil moisture content in gullies (P<0.05) (Table 3). The greater the ability of soil particles to resist runoff scouring damage with more stable soil mechanical properties, which were hard to be destroyed by water flow. Among them, regression analysis on the frictional angle and moisture content of a gully with different activity showed that the internal frictional angle of gully wall and soil moisture content showed a significant linear positive correlation (P<0.05). This indicated that different moisture content

<table>
<thead>
<tr>
<th>Activity characteristics</th>
<th>Gully width /m</th>
<th>Gully depth /m</th>
<th>Gully head-cut height /m</th>
<th>Gully width-depth ratio</th>
<th>Gully wall-slope /°</th>
<th>Gully bed-gradient ratio</th>
</tr>
</thead>
<tbody>
<tr>
<td>Active gully</td>
<td>6.65±1.03a</td>
<td>2.11±0.54a</td>
<td>1.72±0.36a</td>
<td>3.24±0.36b</td>
<td>89.76±0.48a</td>
<td>47.36±5.75a</td>
</tr>
<tr>
<td>Semi-active gully</td>
<td>4.11±0.27b</td>
<td>1.18±0.11b</td>
<td>0.88±0.19b</td>
<td>3.49±0.17b</td>
<td>81.23±0.65a</td>
<td>28.93±2.76b</td>
</tr>
<tr>
<td>Stable gully</td>
<td>3.08±0.42c</td>
<td>0.68±0.11c</td>
<td>0.48±0.1c</td>
<td>4.58±0.26a</td>
<td>60.53±0.99b</td>
<td>15.73±1.89c</td>
</tr>
</tbody>
</table>

Note: Date in the table indicate the mean value ± the standard value. Different lowercase letters for the same stage show significant differences among different gullies at the 0.05 level.

Fig. 3. The shear strength of different parts of the gullies; different lowercase letters (a, b, c) refer to active gully, semi-active gully and stable gully, respectively.
Gully various parts | Gullies with different activity | Regression equation | $R^2$ | $F$ | $P$ |
<table>
<thead>
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<th></th>
</tr>
</thead>
<tbody>
<tr>
<td>Gully head</td>
<td>Stable gully</td>
<td>$y = 26.715x - 0.105$</td>
<td>0.901</td>
<td>27.127</td>
<td>0.014*</td>
</tr>
<tr>
<td></td>
<td>Semi-active gully</td>
<td>$y = 25.822x - 0.179$</td>
<td>0.891</td>
<td>25.839</td>
<td>0.015*</td>
</tr>
<tr>
<td></td>
<td>Active gully</td>
<td>$y = 31.238x - 0.224$</td>
<td>0.899</td>
<td>26.811</td>
<td>0.014*</td>
</tr>
<tr>
<td>Gully wall</td>
<td>Stable gully</td>
<td>$y = 22.516x - 0.122$</td>
<td>0.989</td>
<td>264.48</td>
<td>0.001**</td>
</tr>
<tr>
<td></td>
<td>Semi-active gully</td>
<td>$y = 23.752x - 0.187$</td>
<td>0.944</td>
<td>50.691</td>
<td>0.006**</td>
</tr>
<tr>
<td></td>
<td>Active gully</td>
<td>$y = 19.715x - 0.105$</td>
<td>0.938</td>
<td>45.331</td>
<td>0.007**</td>
</tr>
<tr>
<td>Gully bed</td>
<td>Stable gully</td>
<td>$y = 30.141x - 0.074$</td>
<td>0.862</td>
<td>18.805</td>
<td>0.023*</td>
</tr>
<tr>
<td></td>
<td>Semi-active gully</td>
<td>$y = 38.019x - 0.096$</td>
<td>0.869</td>
<td>19.91</td>
<td>0.021*</td>
</tr>
<tr>
<td></td>
<td>Active gully</td>
<td>$y = 38.335x - 0.159$</td>
<td>0.889</td>
<td>24.101</td>
<td>0.016*</td>
</tr>
</tbody>
</table>

Note: $R^2$, $F$ and $P$ refer to determination, significance test and significance. * and ** indicate the relationship reaches 0.05 and 0.01 significant level, respectively.

Table 3. Regression analysis on the frictional angle and moisture content of a gully with different activeness.

Table 4. Regression analysis on soil cohesion and moisture content of a gully with different activeness.

had a great influence on the internal friction angle. The fitting between friction angle and soil moisture content in the gully bed was the smallest ($R^2<0.9$).

**The Soil Cohesion in Various Parts of a Gully**

Although the moisture content increased as the experiment progressed, soil cohesion showed a power-function decline trend ($P<0.01$) in gully various part and different active gullies (Table 4). It could be seen that cohesion decreased with the increase of soil moisture content, and there showed a significant linear negative correlation ($R^2>0.92$, $P<0.01$). Regression analysis showed that the soil moisture content had a great influence on the soil cohesion in a gully. These results also indicate that soil cohesion was much lower in a stable gully than in the active gully.

Soil mechanical properties and morphological characteristics behave differently under the gully systems. Calculating topographic information entropy showed that the different developmental periods of gully erosion could be determined as active period, semi-active period and active stage, according to different topographic information entropy ($0.1101≤H≤0.1885$, $0.1886≤H≤0.2225$, $0.2226≤H≤0.4000$) (Fig. 4).

The present study has confirmed that the $H$ values of active gully, semi-active gully and stable gully were 0.15, 0.21 and 0.25, respectively. Based on the information entropy classification, the morphological characteristics of gully (gully width, average gully depth, width-depth ratio, bed-gradient ratio, gully wall-slop, and head-cut height) were significantly different. However, it should be noted that gully morphological development was controlled by many factors, such as climate, topography, geotechnical properties, etc., and soil consisted of soil particles, aggregates, soil pores and organic matter, etc. Previous studies had shown that soil characteristics can have a significant impact on the gully erosion process [1, 33]. Moreover, the gully morphological characteristics would be significantly affected by the weathering and soil cracking processes as a result of direct long-term exposure to sunlight [1]. Accordingly, the soil shear strength would also be affected by such processes [33]. It may be worth noting that the soils were largely free.
of stones, as this may affect the soil shear strength processes of various gully parts and different gully activeness.

Field study found that the soil in active gully head was unstable, collapsing frequently and the soil erosion occurred seriously. In conclusion, the morphological characteristics of different active gully (gully width, average gully depth, width-depth ratio, bed-gradient ratio, gully wall-slop, head-cut height) were significantly different. The soil shear strength under different parts of gullies varied with different gully activeness, and the different soil moisture contents significantly affected the internal friction angle and cohesion index of gullies.

In the Dry-Hot Valley region the development of steep headcuts is perceived to be a significant morphological feature of an active gully [3], especially the development of plunge pools at the feet of the headcuts, and several studies have focused on gully headward erosion in this region [30]. Gully morphological development was controlled by many factors, such as climate, topography, geotechnical properties, etc., and the soil consisted of the soil particles, aggregates, soil pores and organic matter, etc. [1-7]. The soil shear strengths under different parts of a gully were varied, and the different soil moisture contents significantly affect the internal friction angle and cohesion index of a gully [15]. In addition to gully development by the influence of water force, uplift force and effective gravity, differences of soil mechanical properties also play a role. The soil with lower cohesion under water erosion was easily loose before causing erosion [1, 5]. Both the soil internal friction angle and soil cohesion were more sensitive to changes in moisture content at low levels than high ones. That could indicate that the effects of soil moisture had a great contribution to gully soils in the Dry-Hot Valley in which the wet and dry seasons were apparently distinct. Studies reported that this result could be ascribed to vertical variation in soil mechanical properties. It should be noted that a significant correlation was found between soil moisture content and soil shear strength [33]. Accordingly, more attention should be paid to vertical variations in soil mechanical properties and weathering processes.

This might be due to the largest gully head-cut height value with undercutting and collapsing activities occurring most frequently [30-33]. In a semi-active gully, the gully head-cut height value was lower than that of an inactive gully, but its gully wall-slope was very steep and more prone to collapse, and thus affected soil stability. The stable gully with the smallest gully bed-gradient ratio was dominated by sediment deposition, and its gully bed terrain tended to be gentle. A field survey found that the growth of gully vegetation with different active degrees of gully in the Dry-Hot Valley were different, and the vertical soil of the gully could be divided into 4 sections with 28 layers. The active gully extremely developed and all parts of that had almost no vegetation growth. The stable gully was barely eroded, and all parts of that almost had vegetation growth. In this paper, the dynamic characteristics of gullies with different active degrees were not analyzed comparatively, and the relevant discussions could be strengthened in the future.

Under the erosion of water flow, the anti-erosion ability of the lower sandy layer was obviously weaker than that of the upper clayey layer, causing it to be washed away first, resulting in the hanging state of the upper clayey layer and creating a collapse potential for the gully head and gully wall. A field study found that soil in an active gully head was unstable, collapsing frequently and soil erosion occurred seriously. Although it would be interesting to assess its morphological characteristics for different gully sections and time periods, the significant advances in the understanding of gully erosion in this region are well worth celebrating; nevertheless, we still do not understand the mechanisms, especially the evolutionary mechanisms for different gully activity. Therefore, further quantitative studies are required to quantify spatial and temporal relationships between morphological characteristics and mechanical properties for different soil types, soil layers and geomorphic locations in the Dry-Hot Valley region.
Conclusions

This research was conducted to determine a gully’s morphological characteristics and its soil mechanical properties in the Dry-Hot Valley region. The results prove that the different developmental periods of gully erosion could be identified as active period, semi-active period and active stage according to different topographic information entropy. Besides, the soil shear strength under different parts of a gully vary with different active degrees of gully. Additionally, the different soil moisture contents significantly affect the internal friction angle and cohesion index of a gully. Furthermore, the gully morphology indicators that affected the gully soil mechanical properties were found to be gully head drop-height, gully wall-slope and bed-gradient ratio. The significant advances in the understanding of gully erosion in this region are well worth celebrating. In particular, the findings of the present study suggest that gully erosion with different topographic information entropy results in different morphological characteristics, and that gully erosion contributes to soil mechanical properties.

Acknowledgements

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Conflict of Interest

The authors declare no conflict of interest.

References

28. SAHANAVIN N., PRUEKSASITT T., TANTRAKARNAPA K. Relationship between PM$_{10}$ and PM$_{2.5}$ levels in high-traffic area determined using path analysis and linear regression. Journal of Environmental Sciences. 69, 105, 2018.